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DISCUSSION OF A NEW METHOD FOR MAPPING IONOSPHERIC CHARACTERISTICS

L. Bossy* and K. Rawer**

*Institut d'Aéronomie Spatiale, B-1180 Bruxelles, Belgium **Albert-Ludwigs-Universität, D-7800 Freiburg, F.R.G.

ABSTRACT

For synthesizing ionospheric maps, it is proposed to use the method of "empirical orthogonal functions" instead of the inversion of the analysis schedule. For interpolating between grid points splining is not the only way. In particular, longitudinal interpolation by a low order Fourier development of the different eigenvectors components appears to be advantageous. Compared with the present synthesis method, important economies in storage capacity and computing time can be achieved.

INTRODUCTION

In 1967, CCIR /1/ adopted a station-based numerical mapping system using a combination of Fourier- and Legendre-analysis (Jones & Gallet, /2/) and a special latitude coordinate called "MODIP" (Rawer, /3/). Data analysis is executed in two steps: first Fourier-analysis station by station of the observed (monthly averaged) diurnal variation; the second step is a world-wide representation of each Fourier coefficient. For synthesis, i.e. in prediction applications, the diurnal variation at an arbitrary position is built-up by the interpolated Fourier-coefficients and this is, of course, also done at the positions of the original stations. This means that basically the same algorithm is used both for analysis and synthesis.

The CCIR-system has two major drawbacks /4/: one is the fact that large areas of the globe, in particular the oceans, have no stations. This is circumvented by an ingenious but criticizable procedure: to fill the gap, data of coastal stations are used once again at a position in the ocean obtained by "suitably shifting" it away from its original position. It was found (CCIR, /1/) that it is best achieved along the Modip coordinate /3/. Another difficulty is due to geophysical facts: the prevailing ionization processes differ over most of the Earth's surface from those over the polar cap. The reason is that there the magnetic fieldlines are open, so that corpuscles can easily arrive from the tail of the magnetosphere.

These difficulties and the analysis method will not be discussed further in the present paper. We will at present consider another method for synthesizing than the inversion of the analysis schedule used. From the analysis, we can start with a global pattern given by a set of input data at locations regularly spaced in longitude and latitude. For convenience we use a prediction map of CCIR to this end.

We apply a synthesis schedule making use of "empirical orthogonal functions". This method was invented by Pearson /5/ and was first applied in ionospheric analysis by Dvinskikh /6/. From the original (rectangular) data matrix, one constructs a square matrix, and determines its eigenvalues and eigenvectors. The approximation is then achieved using eigenvalues in order of decreasing magnitude and their corresponding eigenvectors.

^{*} also Université Catholique de Louvain, B-1348 Louvain-la-Neuve, Belgium.

THE METHOD

A worldmap of foF2 (Figure 1) created by a rectangular matrix C of n rows and m columns; in our example m=24 (longitude) and n=60 (latitude) is used. The rows as well as the columns are independent real vectors; therefore the rank of the matrix is min(m,n). The rows are vectors in an R_n space and the columns are vectors in an R_n space; both spaces are dual. For the presentation of the method, we will admit that m<n.

The first step is to construct an m*m matrix D by (' indicates the transpose) D = C C'(1)

By its construction, D is symmetrical and positive definite; therefore all the eigenvalues of D are real and positive.

Next, we use classical methods to compute the principal components of D by: a) first computing the m eigenvalues λ_i of D and forming an m*m diagonal matrix A by introducing those eigenvalues in descending order of magnitude along the diagonal;

b) second, solving the matrix equation $D U = U \Lambda$

(2) where U is an orthonormal matrix with the eigenvectors e_1, e_2, \ldots, e_m of D as columns. These vectors are known as the principal components of D.

Having determined U (Table 1), the Dvinskikh's method consists in expressing C as the product of U and a new m*n matrix V, such that: C = U V where V = U' C (3) The matrix V (Table 2 shows V') is formed of m row's v_i which are orthogonal n-vectors with decreasing norms. In fact, the square of the norm of each vector v_i is equal to the corresponding eigenvalue (because V V' = A). As can be seen from Table 2 the absolute values of the vector components decrease rather fast.

When, as in our applications, the magnitude of the eigenvalues decreases rapidly, this way of factorizing C is very efficient because of a property described below.

We now call

$$C_{k} = U_{k} V_{k}$$
 with $U_{k} = |e_{1}, e_{2}, \dots, e_{k}, 0, \dots, 0|$ (4)
and $V_{k} = |V_{1}|$
 V_{2}
 V_{k}
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the kth iterate of C.

Since: a) the sum of the squares of the elements of any real matrix A is equal to Since: a) the sum of the squares of the elements of any real matrix A is equal to the trace of the matrix AA', b) this trace is an invariant, which becomes the sum of the eigenvalues when an orthonormal transformation is applied, c) the first k eigenvalues disappear in the trace as a consequence of the orthogonality of the eigenvectors, it follows that the elements of C - Ck, which are the residuals after the kth iteration of C, are such that the sum of their squares is equal to $\lambda_{k+1} + \ldots + \lambda_m$, the sum of the eigenvalues of D between k+1 and m. Therefore, the statistical variance of the residuals after the kth iteration of C equals λ_{k+1} + ... + $\lambda_m.$ In our applications, this variance decreases very rapidly with increasing k.

Thus, the required accuracy will be achieved with a $\ensuremath{\text{pth}}$ iterate, where $\ensuremath{\text{p}}$ is the value of \ensuremath{k} for which the value of the standard deviation

$$\sigma = \sqrt{(\lambda_{p+1} + \ldots + \lambda_p)/N}$$
(5)

Mapping Ionospheric Characteristics

is less than a given precision treshold (N = m n $% \left(N \right) = 0$ is the total number of data). This means then that with an approximation of C $% \left(N \right) = 0$ $C \approx C_{p} = U_{p} V_{p}$ (6)

one can reproduce the grid points of the map with maximum errors of a few σ (the majority lying between -3σ and $+3 \sigma$).

The computation of these N grid values consists in N inner products of p-vectors. These p-vectors are the first p elements in the rows of U which will be called h(LO) (LO longitude) and the first p elements in the columns of V (or in the rows of the transpose) which will be called w(LA) (LA latitude). (Table 3 corresponds to p less than or equal 6)

The element of C at the position LO,LA is given by the inner product C(LO,LA) = (h(LO),w(LA))(7)

Let M = p (m+n) the total number of components of the matrices U_p and V_p , is the quantity of data needed for the representation of the map by the present method. Let Then, the ratio q between the number N = m n of data to be represented, and M the number of data needed

q = N/M = (m n)/[p (m+n)] (8) q can be used as a quality factor. The method saves storage as soon as q is greater than one.

FIRST APPLICATION

As said in the introduction, up to now our numerical analysis has only been carried out for one foF2-worldmap computed with the CCIR coefficients. It corresponds to March 12UT, moderate solar activity.

We thus started with N = 24*60 = 1440 grid-values computed for longitudes between 0 and 345° with 15° steps and latitudes from -88.5 to 88.5° with 3° steps. They constitute the well known matrix C (not reproduced).

Then, the U and V matrices were computed as Tables 1 and 2 (Table 3 was found after reduction with p=6). The units are respectively .001 in Table 1 and .01 Mhz in Table 2. Table 4 shows the statistical distribution of the 1440 residuals displaying for up to 24 iteration steps: a) the variance of the residuals, i.e. the sum of their squares,

b) the maximum of the absolute values of the residuals,

c) the value of the standard deviation,d) the distribution function of the residuals.

We see that if we want to reproduce the 1440 given data with:

high we	precision need p =	9	Emax = .2 MHz iterations and	M = 756	inputs
good we	precision need p =	6	Emax = .5 MHz iterations and	M = 504	inputs
poor we	precision need p =	4	Emax = 1. MHz iterations and	M = 336	inputs

The ratios q = N/M are near 2, 3, and 4 respectively.

The gain achieved by the method is limited by the most rapid changes with position on the map especially in the equatorial region.

Figure 2 shows for p = 6 the geographic regions where the absolute values of the residuals are above .3 MHz. This happens at some 15 grid points, all located in the equatorial region between 30°N and 30°S. It also suggests the reproduction quality one can expect with p = 6.

(11)67

(11)68

INTERPOLATION PROCEDURE

So far, we have only considered the problem of reproducing the original grid points.

In order to compete with existing procedures the method should allow easy interpolation. To this end, one should be able to compute between gridpoints e.g. between the components of each of the p eigenvectors e_i . The same should be possible among the vectors v_i . Spline interpolation is a classical means to this end.

However, we have looked for an analytical approach and tried the following: a) because the longitude domain is normally divided in regular steps, harmonic analysis looks promising. In the present case, for the longitude LO, the kth component of h can be represented by the development

$$h_k(LO) = A_{kO} + \sum_{i=1}^{11} (A_{ki} C_i + B_{ki} S_i) + A_{k12} C_{12}$$
 (9)

where $C_i = \cos(i*L0)$ and $S_i = \sin(i*L0)$ b) because the latitude is only given in the -90°, 90° domain, harmonic development implying symmetry of the poles has been chosen such that for the latitude LA:

$$W_k(LA) = A_{k0} + \sum_{i=1}^{r} A_{ki} C_i$$
 where $C_i = \cos(i*LA)$ (10)

for k equal to 1,...,p. The coefficients occuring in this last development are computed by a least squares method and, by convention, the limiting value r of i is reached when the greatest absolute error is less than .15 MHz. For the evaluation of $h_k(LO)$ only coefficients with i<5 need to be retained.

INTERPOLATION IN THE ORIGINAL MAP

Applying the procedure of interpolation to the 24*60 given data and recomputing the start data using this procedure, we compared the given and recomputed values and found that a good representation (Emax reaches .51 MHz at two locations only) is obtained with 54 coefficients for the h-vectors and 146 coefficients for the w-vectors. So 200 coefficients, with a quality factor q = 7, are good enough to reproduce the grid values as accurately as with the 504 inputs needed by the Dvinskikh's method. The technique also provides an easy way for interpolating the e- and v-vectors.

Interpolating for a position LO,LA gives thus h(LO) and w(LA) and the value of C(LO,LA) at the intended location LO,LA as the inner product C(LO,LA) = (h(LO),w(LA)) (11)

Figure 3 shows the some 30 locations where this procedure leads to absolute errors greater than .3 MHz. They are all located between $30^{\circ}N$ and $30^{\circ}S$, and show no distinct structure. Emax only increases insignificantly (from .49 to .51 MHz) and there are a few more large error locations. In fact, almost all residuals fall into the noise level so that the proposed interpolation procedure causes no significant degradation of the results with the original Dvinskikh method.

FURTHER DEVELOPMENTS

Some numerical experiments were made to try to optimize the interpolation procedure.

First, we considered the effect of (regular) latitudinal spacings with 6, 9 and 12°. With p = 6, the 1440 values at the grid points were computed and compared with the original ones.

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Statistical comparison gave the following results :

	Emax	Number of points with residuals > .3 MHz	Number of storages M	Value of q = N/M
<u>-6*</u>	.46 MHz	29	193	7
9•	.56 MHz	67	174	8
12.	>2. MHz	from 60°N to 60°S		

It seems that about 10° is the maximum regular spacing permitted in latitude. With 9°, a 10% economy of storage is reached, at the cost of some loss of precision and of some spreading of the residuals towards the temperate latitudes (Figure 4). Obviously a 12° spacing cannot be satisfactory because it is unable to reproduce the equatorial and surrounding regions.

In a second series of experiments, tests with non regular spacings have been made. In particular, we have introduced a narrower spacing in the equatorial region where the variability is most important. For three schemes, the results are :

6° general spacing with 3° near the equator Emax = .53 MHz 33 locations 193 storages q = 79° general spacing with 3° near the equator Emax = .60 MHz 80 locations 186 storages q = 89° general spacing with 6° near the equator (Figure 5) Emax = .51 MHz 38 locations 186 storages q = 8

We note that the non-regular 9° and 6° scheme is almost as good as the regular with 6° one. The more economic regular 9° scheme is less precise and shows a substantially greater number of residuals above .3 MHz. The unsatisfying result obtained with the 9° and 3° may be due to the perturbation introduced by higher orders in the Fourier decompositions; these are needed in order to reproduce the strong variations in the equatorial zone but reappear as undulating perturbations in the temperate latitudes.

CONCLUSIONS

It was the goal of this study to determine, by numerical experiments, the storage economy obtainable when establishing world maps by the Dvinskikh's method or some extension of it.

With our method of analytical interpolation, the 1440 grid points used for the definition of the map could satisfyingly be recomputed with only 175 to 200 storages, thus with a quality factor between 7 and 8. At a few locations, in the equatorial region, the residuals went up to \approx .5 MHz. Our procedure also allows to interpolate in the map with no significant loss of precision. The original Dvinskikh's method gives a representation which is slightly better but needs follow-up local interpolations. Some 500 storages are necessary, leading to a quality factor of only 3. Thus, the introduction of descriptive functions in place of the original discrete values of the elements of the e- and v-vectors is advantageous.

There is, of course, no unique way of looking for such interpolation functions. They have to be chosen having in mind the type of map considered : worldmap or regional map.

The good results reported have perhaps been influenced by the fact that the grid data were derived from a smooth Legendre expansion without noise.

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ACKNOWLEDGMENTS

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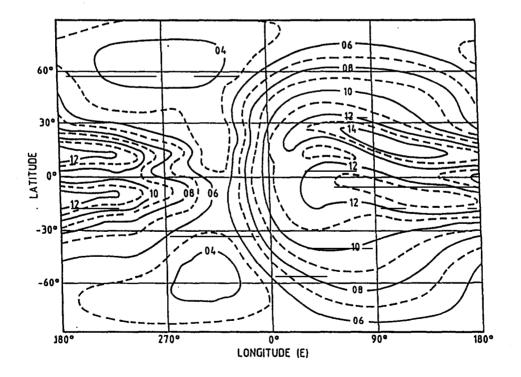


Fig. 1. Worldmap of foF2.

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236	-140	-24	-10	26	-7	79	17	-283	272	-165	-45	95	112	-53	-196	20	167	-291	446	-411	333	-255	-33
134	-1/3	214	55	-189	-12	121	32	-185	-70	297	-221	249	-300	21	155	265	-476	196	-34	-208	191	164	-228
104	- <u>-</u>	-32	-20	44	2	0	S	38	-119	75	102	-394	605	-461	224	-226	209	-69	-31	157	-172	71	89
62	13	-38	-20	36	12	-70	156	-259	348	-287	178	-228	226	66	-379	363	-269	242	-119	121	-219	211	-148
152	4-	-122	-29	116	-16	14	-93	310	-474	323	-85	-136	236	119	-511	346	-75	-10	111	-55	34	-94	-54
232	-251	12	-23	20	0	170	-173	160	43	-254	-105	294	166	-229	-102	-76	- 59	493	-234	72	105	-432	171
1	32	-95	-115	207	24	-288	512	-396	-54	318	-336	213	37	-45	-97	-39	136	27	-167	223	-60	-183	128
ω	-29	155	- 144	55	54	-277	318	33	-228	-114	437	-313	-88	32	222	65	-294	33	0	-10	326	-372	131
399	-164	23	13	-17	-15	141	-238	-205	109	193	124	-230	-121	381	-33	-177	182	-189	-360	309	152	-167	-221
370	-277	146	-216	110	66	-149	188	194	-226	-277	156	312	-216	-70	-29	-22	124	-142	152	167	-279	187	308
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IABLE 1 Elements of Matrix U (in .001).

Mapping Ionospheric Characteristics

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0 -141 69 4 37 4 -12 -13 -8 2 0 0 0 25 -152 58 -15 28 -2 -5 -12 -8 0 0 0 0	25 -152 58 -15 28 -2 -5 -12 -8 0 0 0 45 -156 43 -32 17 -8 1 -10 -5 -1 0 -1 0 54 -153 27 -44 8 -14 9 -6 -3 -3 0 0 0 52 -145 10 -49 0 -19 17 -4 0 -6 0 0 0 33 -131 -4 -49 -7 -23 22 -1 3 -8 0 0 0 0 -113 -18 -42 -12 -25 24 0 5 -9 0 3 0 -45 -91 -30 -32 -17 -24 26 2 8 -10 0 4 0	25 -152 58 -15 28 -2 -5 -12 -8 0 0 0 45 -156 43 -32 17 -8 1 -10 -5 -1 0 -1 0 54 -153 27 -44 8 -14 9 -6 -3 -3 0 0 0 52 -145 10 -49 0 -19 17 -4 0 -6 0 0 0 33 -131 -4 -49 -7 -23 22 -1 3 -8 0 0 0 0 -113 -18 -42 -12 -25 24 0 5 -9 0 3 0
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 $\underline{\text{TABLE 2}}$ Elements of the Transpose of Matrix V (9 last columns omitted).

Mapping Ionospheric Characteristics

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Ī	217	97	21	55	312	497
ł	236	170	80	41	228	358
۱	245	239	9	178	229	-138
l	249	278	- 50	265	215	-351
۱	251	302	-18	284	-25	-339
ł	252	279	67	140	-383	-72
۱	252	185	161	-112	-520	187
ł	250	75	186	-249	-269	128
l	244	-2	174	-274	-17	-23
ł	235	- 60	164	-285	87	-102
	225	-122	150	-283	131	-123
	211	-193	111	-229	207	-165
	198	-255	64	-115	203	-209
	187	-291	44	9	27	-151
	182	-312	61	154	-176	34
	179	-338	49	314	-111	81
ł	170	-309	-42	349	-16	36
	154	-222	-147	208	-75	90
1	142	-156	-213	79	-118	102
	131	-111	-306	-25	-112	-2
	114	-25	-466	-171	-115	-110
	112	57	-525	-261	-49	-88
	141	80	-376	-142	72	132
	183	74	-146	26	226	350

2617129 -344 -65 -46 132650-63-284-73-58-72685-1-223-72-69-29272552-163-62-79-47277496-108-42-87-602832131-60-16-91-672900158-219-92-682978178739-90-6330691942468-82-5131682043094-70-53327721125115-54-9339321411130-3317317211-8137-8443645202-3213617683780185-5712745843922156-7810872914074114-93829785424150-9548117664429-37-8311130304643-159-54-23134-184886-312-10-46122-725145-48438-4992-1175389-63172-2542-1345551-6556723-22-1065562-5771283-90-375398<	2559	-265 -	471	÷ ·	-28	- 1	
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	2431	-279	-447	- 33	-28	47	1

<u>TABLE 3</u> Matrices U and Transposed V as Used for Numerical Experiments (U in .001 and V in .01 MHz).

L. Bossy and K. Rawer

							_							-
р	Variance	Em	σ		Frequencies									
01	1145.8253	303	892	0	0	12	87	302	617	314	98	10	0	0
02	348.9192	251	492	0	2	14	33	347	667	299	45	16	14	0
03	143.8461	156	316	1	3	9	71	370	558	333	85	8	2	0
04	75.7184	119	229	1	3	14	77	272	678	317	56	18	22222	2
05	37.9637	76	162	1	6	10	82	276	661	321	66	14	2	1
06	18.8878	49	115	0	2	23	69	302	648	302	83	9	2	0
07	11.3185	41	89	1	3	19	62	304	661	292	82	14	2	0
08	5.5777	40	62	4	3	17	64	252	772	256	51	12	6 6 7	3
09	1.7410	20	35	3	7	14	42	268	785	255	39	15	6	6
10	.7720	17	23	3	1	15	59	307	680	315	40	10		3
11	.3958	16	17	3 3 3	1	6	66	284	710	309	47	6	4	4
12	.2277	15	13		2	8	40	269	797	281	29	6	0	б
13	.1203	8	9	0	4	17	58	300	673	308	67	10	1	2
14	.0636	3	7	0	2	14	75	278	682	293	77	16	2	1
15	.0275	3	4	1	6	11	64	264	752	250	72	15	2 2 2 3	3
16	.0103	2	3	2	3	13	64	270	717	290	66	11	2	2
17	.0046	1	2	0	5	14	64	301	680	293	67	12	3	1
18	.0027	1	1	1	7	15	58	279	732	253	74	17	4	0
19	.0016	1	1	0	4	19	68	250	759	254	64	17	5	0
20	.0013	0	1	1	2	19	70	267	723	266	74	14	5 3 5 6	1
21	.0009	0	1	2	4	17	61	255	757	259	62	16	5	2
22	.0006	0	1	2	5	15	66	240	779	248	60	18	6	1
23	.0005	0	1	1	4	22	53	270	737	281	47	15	9	1
24	.0002	0	0	1	6	16	55	268	745	268	58	16	5	2

<u>TABLE 4</u> p: order of the iteration. Variance: sum of the squares of the residuals (in MHz²). Column \mathcal{E}_m : maximum absolute value amongst the residuals (in .01 MHz). Column σ : standard deviation in (.001 MHz). Last 11 columns: statistical distribution of the deviations central column around zero (i.e. in $-\frac{1}{2}\sigma$ and $\frac{1}{2}\sigma$), side columns for larger deviations (in 1. σ steps).

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-14	Fig. 2.		100 (E) 100 (E)) 4 · · · · · · · · · · · · · · · · · ·		r n Fig. 3.	,, VOC 801	ų. u	14.

Fig. 2. Geographical distribution of the locations giving absolute errors greater than .3 MHz (dots identify error peaks) with Dvinskikh's original method (latitudinal spacing 3°, 540 coefficients), for p = 6.

<u>Fig. 3.</u> Same as Figure 2 when the Fourier interpolation method (with 200 coefficients) is applied (p = 6).

